

Buoyancy and Entropy

Specific Volume: $\alpha = 1/\rho$

Specific Entropy: s

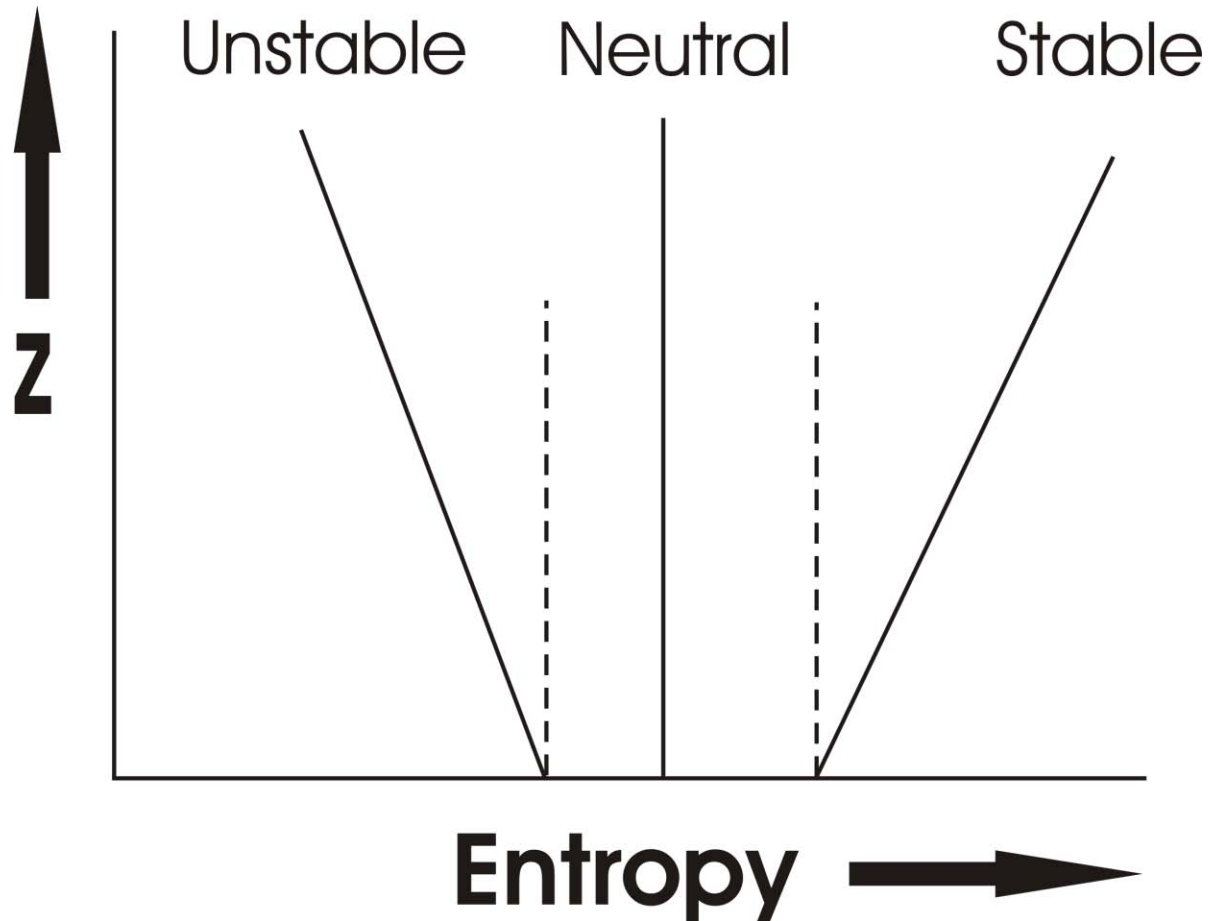
$$\alpha = \alpha(,$$

$$(\delta\alpha)_p = \left(\frac{\partial\alpha}{\partial s}\right)_p \delta s = \left(\frac{\partial T}{\partial p}\right)_s \delta s$$

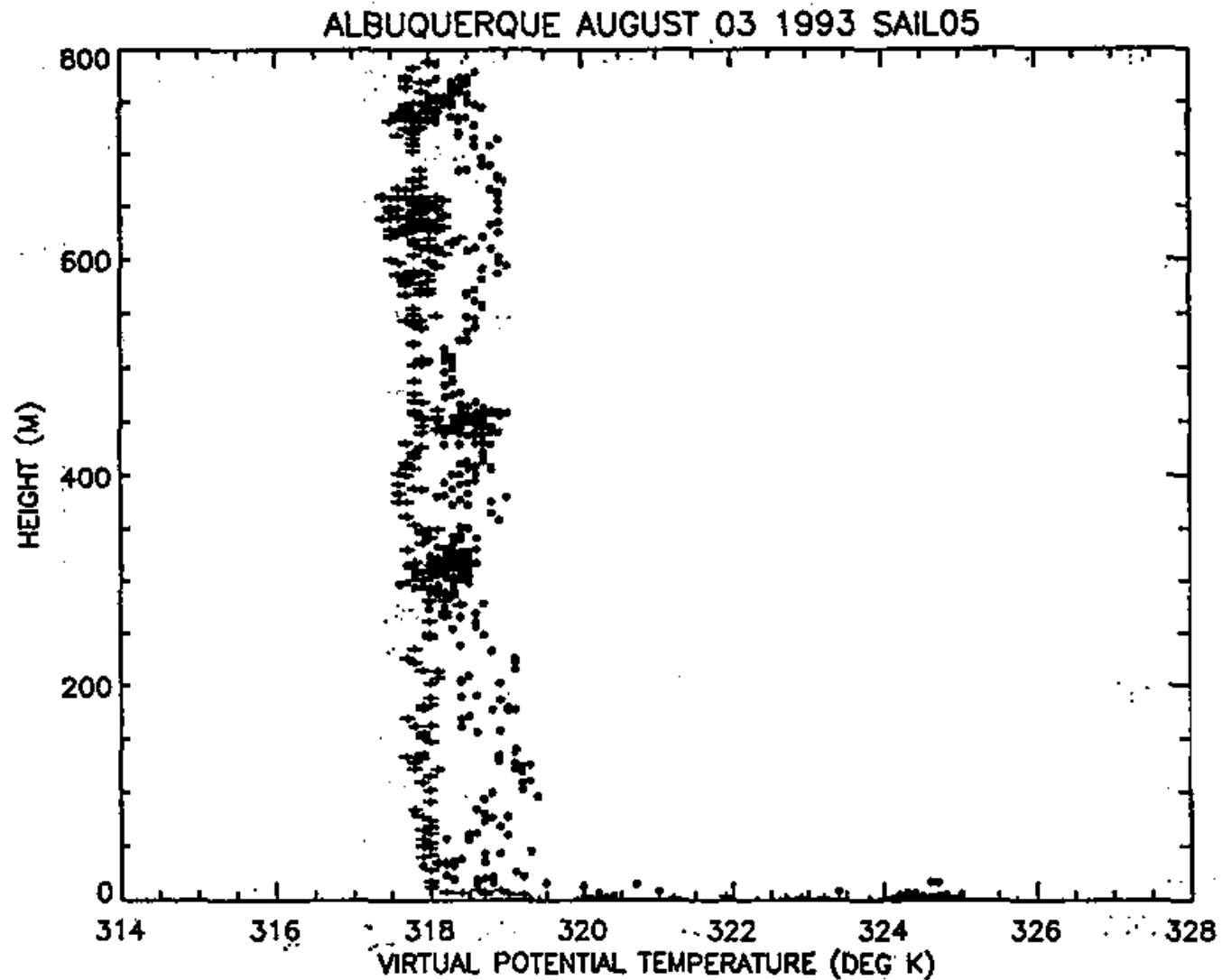
$$B = g \frac{(\delta\alpha)_p}{\alpha} = \frac{g}{\alpha} \left(\frac{\partial T}{\partial p}\right)_s \delta s = -\left(\frac{\partial T}{\partial z}\right)_s \delta s \equiv \Gamma \delta s$$

Note: For ideal gas: $\Gamma = g / c_p$

Earth's atmosphere: $\Gamma = 1 K / 100 m$



Model Aircraft Measurements (Renno and Williams, 1995)



Water Variables

Mass concentration of water vapor (*specific humidity*):

$$q \equiv \frac{M_{H_2O}}{M_{air}}$$

Vapor pressure (partial pressure of water vapor): e

Saturation vapor pressure: e^*

C-C:
$$e^* = 6.112 \text{ hPa} e^{\frac{17.67(T-273)}{T+30}}$$

Relative Humidity:
$$H \equiv \frac{e}{e^*}$$

The Saturation Specific Humidity

Ideal Gas Law:

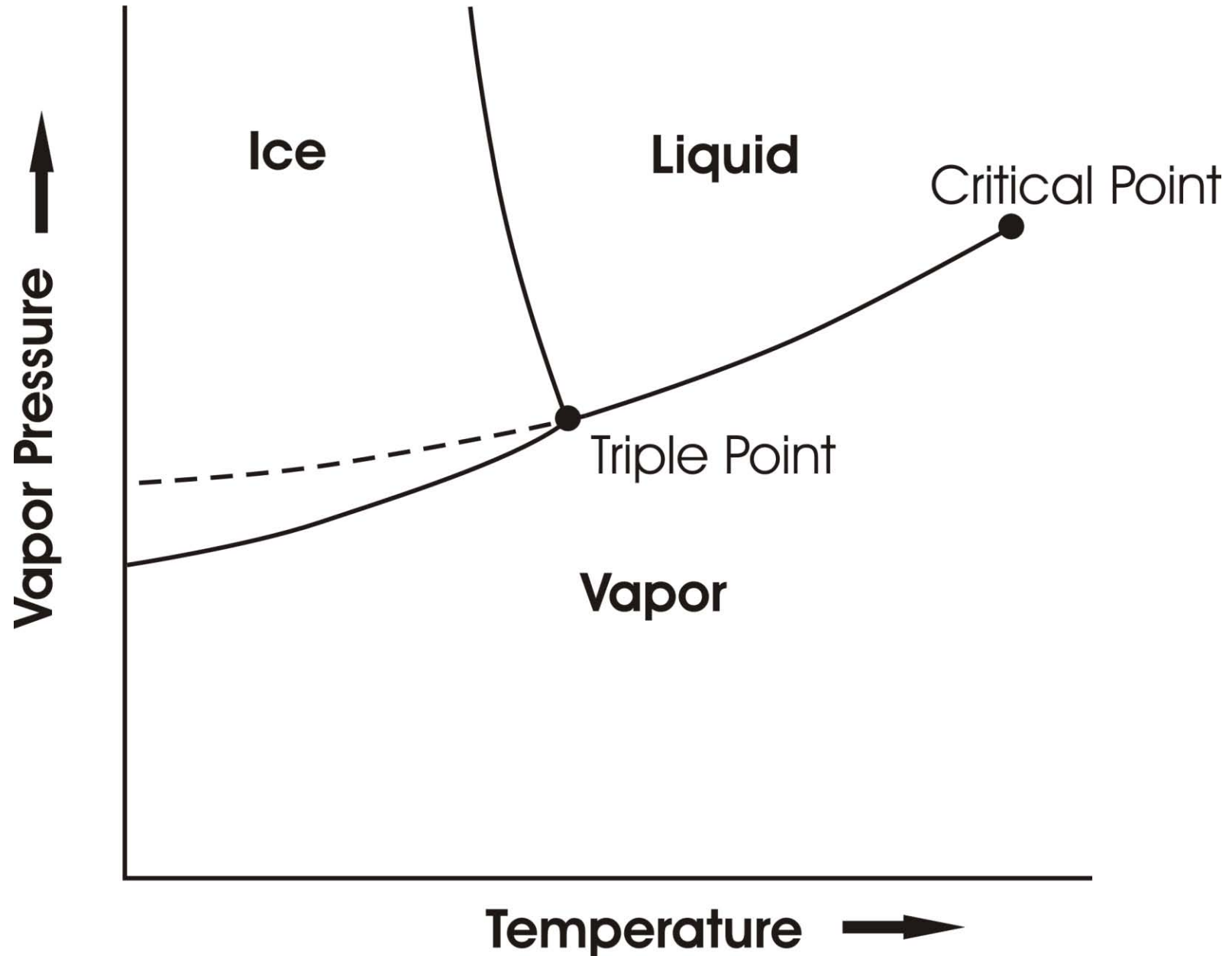
$$p = \rho \frac{R^* T}{\bar{m}}$$

$$e = \rho_v \frac{R^* T}{m_v}$$

$$q = \rho_v / \rho = \frac{m_v}{\bar{m}} \frac{e}{p}$$

$$q^* = \frac{m_v}{\bar{m}} \frac{e^*}{p}$$

Phase Equilibria



Bringing Air to Saturation:

$$e = qp \left(\frac{\bar{m}}{m_v} \right)$$

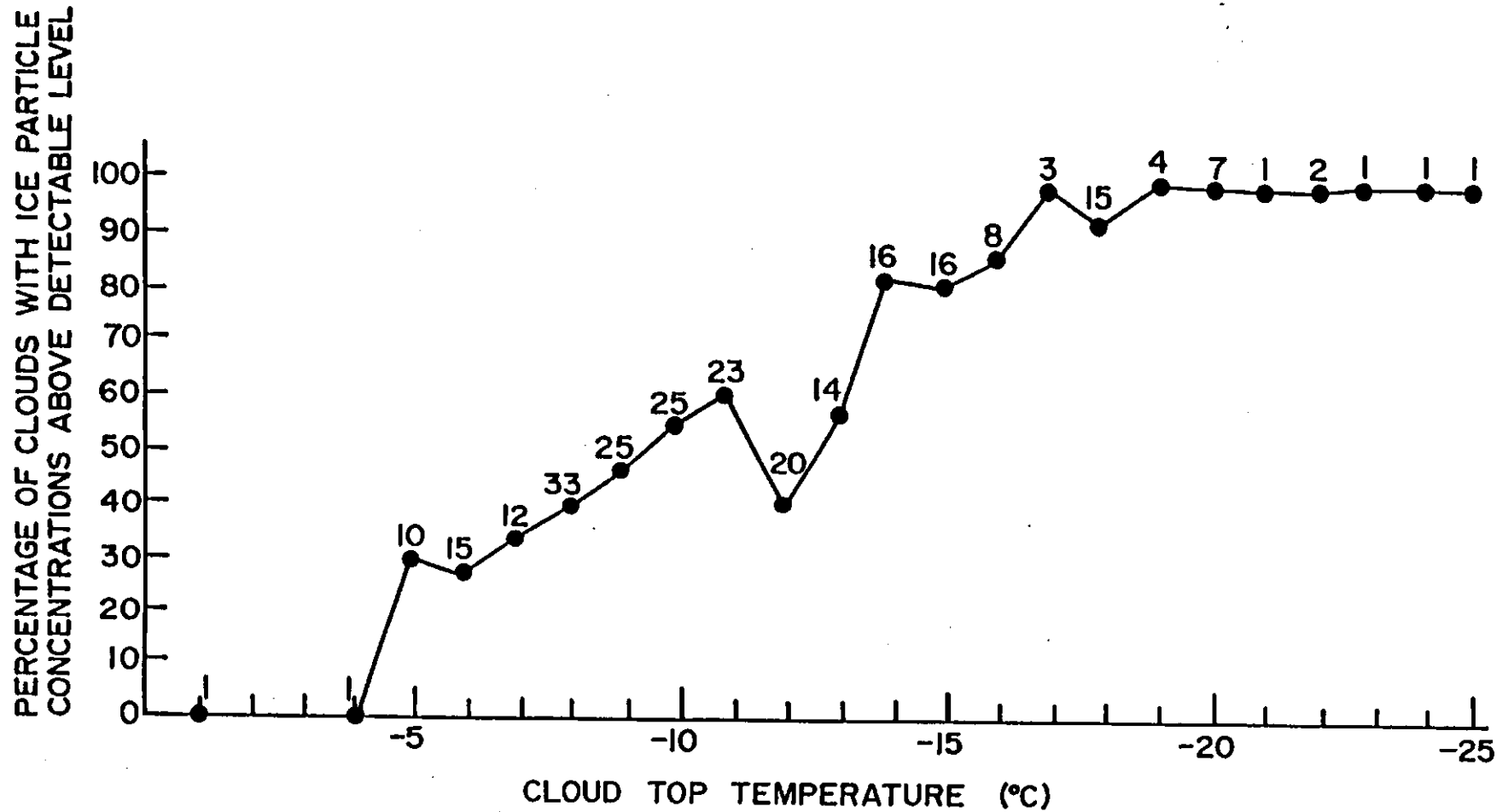
$$e^* = e^*(T)$$

1. Increase q
2. Decrease $e^*(T)$

When Saturation Occurs...

- Heterogeneous Nucleation
- Supersaturations very small in atmosphere
- Drop size distribution sensitive to size distribution of cloud condensation nuclei

Ice Nucleation Problematic



Precipitation Formation:

- Stochastic coalescence (sensitive to drop size distributions)
- Bergeron-Findeisen Process
- Strongly nonlinear function of cloud water concentration
- Time scale of precipitation formation ~10-30 minutes

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